

This article was downloaded by:

On: 25 January 2011

Access details: *Access Details: Free Access*

Publisher *Taylor & Francis*

Informa Ltd Registered in England and Wales Registered Number: 1072954 Registered office: Mortimer House, 37-41 Mortimer Street, London W1T 3JH, UK



## Separation Science and Technology

Publication details, including instructions for authors and subscription information:

<http://www.informaworld.com/smpp/title~content=t713708471>

## Selective Separation of Magnesium from Olivine Minerals

S. E. Ghazy<sup>a</sup>; M. A. Kabil<sup>a</sup>; A. M. Abeidu<sup>a</sup>; N. M. El-Metwally<sup>a</sup>

<sup>a</sup> DEPARTMENT OF CHEMISTRY FACULTY OF SCIENCE, MANSOURA UNIVERSITY, MANSOURA, EGYPT

**To cite this Article** Ghazy, S. E. , Kabil, M. A. , Abeidu, A. M. and El-Metwally, N. M.(1996) 'Selective Separation of Magnesium from Olivine Minerals', Separation Science and Technology, 31: 6, 829 — 838

**To link to this Article:** DOI: 10.1080/01496399608001327

**URL:** <http://dx.doi.org/10.1080/01496399608001327>

PLEASE SCROLL DOWN FOR ARTICLE

Full terms and conditions of use: <http://www.informaworld.com/terms-and-conditions-of-access.pdf>

This article may be used for research, teaching and private study purposes. Any substantial or systematic reproduction, re-distribution, re-selling, loan or sub-licensing, systematic supply or distribution in any form to anyone is expressly forbidden.

The publisher does not give any warranty express or implied or make any representation that the contents will be complete or accurate or up to date. The accuracy of any instructions, formulae and drug doses should be independently verified with primary sources. The publisher shall not be liable for any loss, actions, claims, proceedings, demand or costs or damages whatsoever or howsoever caused arising directly or indirectly in connection with or arising out of the use of this material.

## Selective Separation of Magnesium from Olivine Minerals

---

S. E. GHAZY, M. A. KABIL, A. M. ABEIDU,  
and N. M. EL-METWALLY

DEPARTMENT OF CHEMISTRY

FACULTY OF SCIENCE

MANSOURA UNIVERSITY

P.O. BOX 66, MANSOURA, EGYPT

### ABSTRACT

The aim of this investigation was the selective separation of Mg(II) from olivine, (Mg,Fe)SiO<sub>4</sub>. The feasibility of leaching of Mg<sup>2+</sup>, Fe<sup>2+</sup>, and Fe<sup>3+</sup> from olivine and serpentinized olivine with HCl or H<sub>2</sub>SO<sub>4</sub> solutions of pH ~ 2 or 3.5 has been examined. The results obtained showed that, under identical conditions, HCl solution is more effective than H<sub>2</sub>SO<sub>4</sub> in leaching of Mg<sup>2+</sup> as well as Fe<sup>2+</sup>, and that the optimum pH is <3.5. Magnesium ions are selectively and quite effectively separated from the leaching solution by precipitation of Mg(OH)<sub>2</sub> at pH ~ 11.5 since most, if not all, other metal ions that exist in natural samples are pre-precipitated as hydroxides and separated by filtration. Alternatively, Mg<sup>2+</sup> is selectively separated from the leaching solution by flotation as magnesium oleate in the 7.0–8.5 pH range.

### INTRODUCTION

Magnesium is the lightest of the industrial metals employed on a large scale (1–4). It finds extensive uses in bombs, pyrotechnics, photographs, cameras, and a number of household gadgets; as a reducing agent for the production of Zr, Hf, U, Be, etc.; as cathodic protection for steel pipelines; and as a negative electrode in primary electric cells. Moreover, its alloys are used in the manufacture of many machinery parts for various industries: automotive, textile, and aircraft. Hence, the recovery of magnesium from its natural sources is a vital process.

Magnesium does not occur naturally as a native element but always in the combined state in the form of such minerals and salts as magnesite,  $\text{MgCO}_3$ ; dolomite,  $\text{CaMg}(\text{CO}_3)_2$ ; talc,  $\text{Mg}_3\text{SiO}_4\text{O}_{10}(\text{OH})_2$ ; brucite,  $\text{Mg}(\text{OH})_2$ ; chrysotile,  $\text{Mg}(\text{Si}_2\text{O}_5)(\text{OH})_4$ ; tremolite,  $\text{Ca}_2\text{Mg}_5\text{Si}_8\text{O}_{22}(\text{OH})_2$ ; epsomite,  $\text{MgSO}_4 \cdot 7\text{H}_2\text{O}$ ; carnallite,  $\text{MgCl}_2 \cdot \text{KCl} \cdot 6\text{H}_2\text{O}$ ; olivine,  $(\text{Mg},\text{Fe})\text{-SiO}_4$ ; and serpentine,  $(\text{Mg},\text{Fe})_3\text{Si}_2\text{O}_5(\text{OH})_4$  (5, 6). Of these, magnesite and dolomite are presently considered to be of commercial importance. This may due to their occurrence in vast quantities coupled with the ease of extraction of magnesium from them. However, no data are available in the literature about the extraction of magnesium from silicate minerals. Therefore, the objective of this investigation is to study the feasibility and amenability of selective separation of  $\text{Mg}^{2+}$  from olivine and serpentinized olivine. This goal is achieved by leaching with HCl at  $\text{pH} < 3.5$ . Then the pH of the leachate is increased to an extent permitting the precipitation of most concomitant species (especially  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$ ) with magnesium, followed by filtration. Magnesium ions in the filtrate are collected in the pure state by either precipitation with lime or flotation using the very cheap and available oleic acid soap as a collector after adjusting the pH to the optimum.

## EXPERIMENTAL

### Reagents

All chemical reagents used were of analytical or equivalent grades. Oleic acid used for the preparation of sodium oleate was of technical grade.

### Minerals

Olivine minerals consist of transparent round pebbles with a green or yellowish-green color. Distributed rocks of magnesium minerals occur in the Eastern Desert of Egypt (7). Reasonably pure crystals of olivine and serpentinized olivine were used throughout this investigation. Infrared analyses showed that the samples used were clean and free from locked particles and gangue. They were crushed, ground in an agate mortar, and screened to a  $-100 + 200$  mesh fraction for leaching tests and to a  $-300$  mesh fraction for infrared studies.

### Apparatus

A Perkin-Elmer 2380 atomic absorption spectrophotometer was used throughout for quantitative determination of metal ions. The infrared spectra were recorded on a Perkin-Elmer 1430 ration spectrophotometer. The pH measurements were carried out using a Hanna Instruments 8519 digital pH meter.

## Procedure

### *Leaching of $Mg^{2+}$ and $Fe^{2+}$*

The feasibility of leaching  $Mg^{2+}$  and  $Fe^{2+}$  from the surface of olivine particles was studied as follows: 5 g of an olivine – 100 + 200 mesh fraction was introduced into a 250-mL beaker. To this olivine sample, 100 mL of aqueous hydrochloric or sulfuric acid solution of the desired pH was added. The pulp was stirred by means of a magnetic stirrer for the required time. The leaching rate (dissolution) of  $Mg^{2+}$  and  $Fe^{2+}$  and/or  $Fe^{3+}$  was determined by periodically pipetting a 25-ml aliquot from the supernatant solution for quantitative determination of  $Mg^{2+}$  and  $Fe^{2+}$  ( $Fe^{3+}$ ) as described previously (8–10).

The selective separation of  $Mg^{2+}$  from the leachate was achieved either by precipitation of  $Mg^{2+}$  as  $Mg(OH)_2$  at pH  $\sim$  11.5 after filtration of the hydroxides of the contaminated metals or by ionic flotation of it as magnesium oleate.

### *Ionic Flotation*

Ionic flotation of  $Mg^{2+}$  from the leaching solution of olivine or serpentized olivine was carried out by precipitating metal hydroxides (especially that of iron) and filtering. To 50 mL of the filtrate (2.5 g/L  $Mg^{2+}$ ) in a 100 mL beaker, 12 mL of sodium oleate (1 g/L) was added and the pH was adjusted to the optimum pH range (7.5–8.5). The magnesium oleate formed is self-flatable and does not need any air bubbles. The optimum time of flotation was 5 minutes, after which the floated magnesium oleate was separated from the mother liquor by means of separating funnels. The recovery and effectiveness of selective separation were calculated by quantitative determination of  $Mg^{2+}$  in the residual solution from the relation,  $R = (C_i - C_r)/C_i \times 100\%$ , where  $C_i$  and  $C_r$  denote the initial and residual concentrations of  $Mg^{2+}$  in the mother liquor as described previously (11). Alternatively, the recovery may be calculated by quantitative determination of  $Mg^{2+}$  in the float after dissolving it in a mixture of concentrated  $HNO_3$  and methyl alcohol (1:3).

## RESULTS AND DISCUSSION

### *Leaching of $Mg^{2+}$ and $Fe^{2+}$*

Figures 1 and 2 report the feasibility of leaching of  $Mg^{2+}$ ,  $Fe^{2+}$ , and  $Fe^{3+}$  from olivine and serpentized olivine with HCl and  $H_2SO_4$  solutions of pH 2 and 3.5 at different times. Note that there is a direct and close relationship between the amount of  $Mg^{2+}$  and that of  $Fe^{2+}$  leached, but the rate of dissolution is not constant. The rate of dissolution is relatively

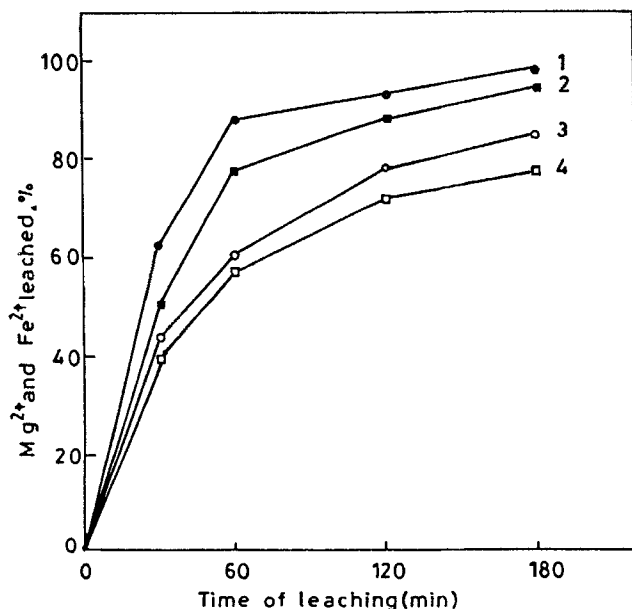


FIG. 1 Leaching of  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  from olivine ( $-100+200$  mesh fraction) by HCl as a function of time and pH. Curves 1 and 3:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  at pH 2, respectively. Curves 2 and 4:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  at pH 3.5, respectively.

high during the initial 60 minutes, after which it increases regularly up to 2 hours. Any further increase of the time of leaching has only a slight effect.

An increase of pH of the leaching solutions from 2 to 3.5 caused a substantial change in the rate and in the total amount of leachable  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ , and  $\text{Fe}^{3+}$ . An interpretation of these results may be possible if we take into consideration the chemical characteristics of olivine- and serpentine-forming cations, namely  $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ , and  $\text{Fe}^{3+}$  in acid medium. Since magnesium only exists in one valence state (I), whereas iron is capable of existing in more than one valence state in combination with anions, there is a difference in the solubilities between the valence states of iron in acid medium, and the prevailing redox potential may seriously affect the mobility of iron. Indeed, it has been well established that precipitation of  $\text{Fe}(\text{OH})_3$  occurs in the pH range 2.1–3.5 whereas precipitation of  $\text{Fe}^{2+}$  as  $\text{Fe}(\text{OH})_2$  requires a much higher pH value (pH 7.5–9.5) (12, 13).

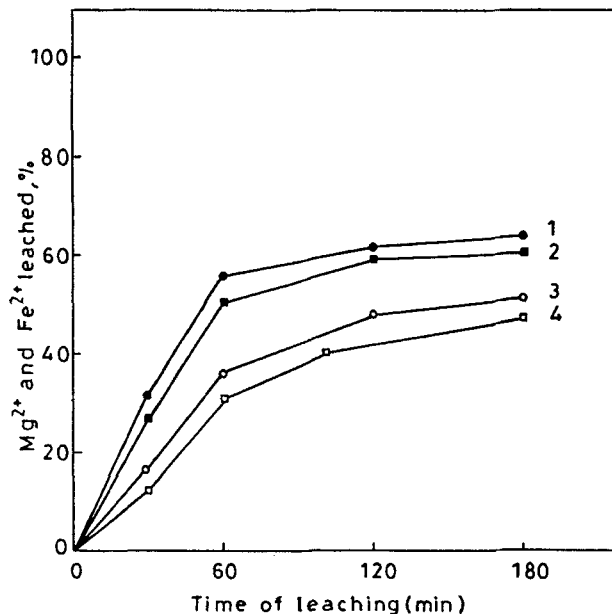
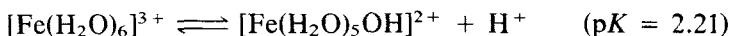


FIG. 2 Leaching of  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  from olivine ( $-100+200$  mesh fraction) by  $\text{H}_2\text{SO}_4$  as a function of time and pH. Curves 1 and 3:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  at pH 2, respectively. Curves 2 and 4:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  at pH 3.5, respectively.

It seems reasonable to assume that at  $\text{pH} \leq 3.5$ , olivine and serpentine break down to the individual cations ( $\text{Mg}^{2+}$ ,  $\text{Fe}^{2+}$ , and  $\text{Fe}^{3+}$ ) contained therein, which would be expected to assume their preferential octahedral structures:  $[\text{Mg}(\text{H}_2\text{O})_6]^{2+}$ ,  $[\text{Fe}(\text{H}_2\text{O})_6]^{2+}$ , and  $[\text{Fe}(\text{H}_2\text{O})_6]^{3+}$ . This is in accord with the well-established phenomenon that dissolution of minerals is enhanced by reagents able to complex minerals which form cations (14, 15). This may also explain the sensitivity of serpentine, which contains  $\text{Fe}^{3+}$ , to the pH values of the leaching solutions (Figs. 3 and 4). Thus,  $[\text{Fe}(\text{H}_2\text{O})_6]^{3+}$  is a fairly strong acid and undergoes acid ionization such as (16):



The first ionization constant for  $[\text{Fe}(\text{H}_2\text{O})_6]^{3+}$  is comparable to the second ionization constant for sulfuric acid, being about  $2.1 \times 10^{-2}$ .

Accordingly, when conditions are suitably reducing and the pH is  $\leq 3.5$ , much of the iron may be leached from olivine, whereas if the environment

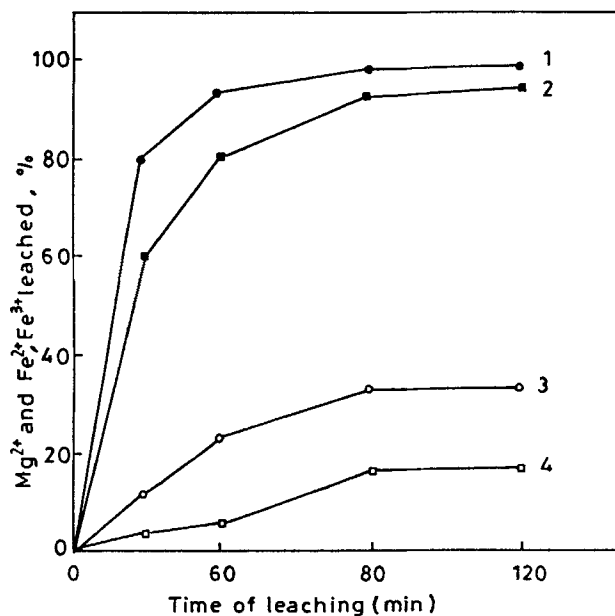


FIG. 3 Leaching of  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}(\text{Fe}^{3+})$  from serpentinized olivine ( $-100+200$  mesh fraction) by HCl as a function of time and pH. Curves 1 and 3:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}(\text{Fe}^{3+})$  at pH 2, respectively. Curves 2 and 4:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}(\text{Fe}^{3+})$  at pH 3.5, respectively.

is oxidizing and the pH is  $\geq 3.5$ , iron is rendered immobile and is stabilized in the Fe(III) state. Moreover, it has been reported that solutions of salts of trivalent metal ions are acidic due to their hydrolysis. Only the formation of mononuclear hydrolysis products was considered for a long time, but Bjerrum and others assumed that in the hydrolysis of some metal(III) cations, polynuclear complexes, e.g.,  $\text{M}_2(\text{OH})_2^{4+}$  and  $\text{M}_2(\text{OH})_4^{2+}$ , are also formed. In the last two decades the hydrolysis of different metal ions has been extensively investigated (15). Moreover, it is also well established that metal cations exist in acidic solution as aquo complexes and that different mono- and polynuclear oxo- and hydroxo-complexes are formed on adding alkali.

The role played by hydrogen ion in leaching cannot be ignored. Not only does it initiate distribution of the crystal structure but its concentration also influences the solution and precipitation of some of the released ions. However, hydrogen ion concentration is not an independent variable in leaching but rather is a function of several interrelated factors: composi-

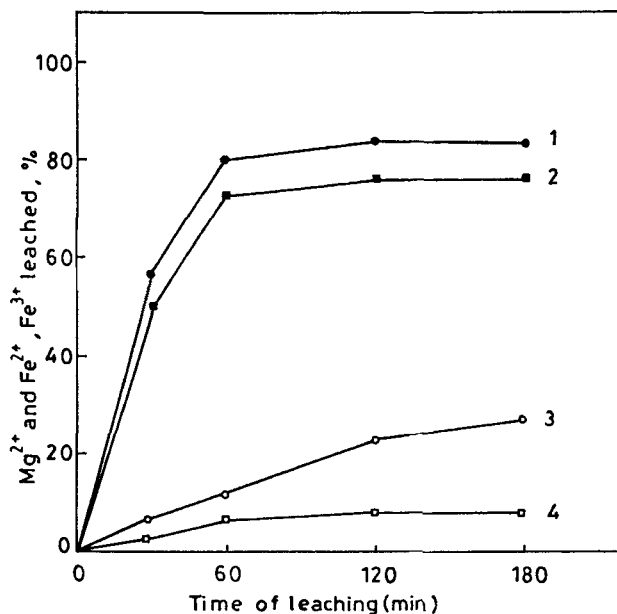


FIG. 4 Leaching of  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  ( $\text{Fe}^{3+}$ ) from serpentinized olivine ( $-100+200$  fraction) by  $\text{H}_2\text{SO}_4$  as a function of time and pH. Curves 1 and 3:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  ( $\text{Fe}^{3+}$ ) at pH 2, respectively. Curves 2 and 4:  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  ( $\text{Fe}^{3+}$ ) at pH 3.5, respectively.

tion and structure of the minerals to be leached, the rate of leaching, the nature and, in particular, the cation-exchange capacity of the residual mineral product. Minerals break down primarily because some of their constituent atoms and ions are dissolved and effectively removed from the environment. The loss of those constituents renders the existing mineral structure unstable, and new crystalline phases tend to form. Thus, when conditions are suitable for reduction and the  $\text{pH} < 3.5$ , iron may be leached from iron-bearing minerals such as olivine and serpentine. In contrast, if the environment is suitable for oxidation, iron is stabilized even in a fairly strong acid media ( $\text{pH} < 3.5$ ) (Figs. 1–4).

It has been reported that the dissolution and behavior of various ion-leaching reactions can be related directly to their ionic potential, which is a fundamental property of the element because it is closely related to the element's electronegativity. Ionic potential is expressed numerically as the ratio of the charge in valency units to the ionic radius in Angstrom units. Consequently, both  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$  have an ionic potential of 2.6



(14), i.e., less than 3, making them tend to pass to the leaching solution in ionic form (14). On the other hand,  $\text{Fe}^{3+}$  has an ionic potential of 4.5, i.e., much greater than 3 and less than 9, which makes it precipitate and become concentrated in the residue (14).

Figures 1–4 indicate the superiority of HCl over  $\text{H}_2\text{SO}_4$  in leaching olivine and serpentinized olivine. This may be attributed to its nonoxidizing properties coupled with the ability of  $\text{Cl}^-$  ions to form complex soluble anions with olivine-forming cations, namely  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$ , of the type  $\text{MgCl}_4^{2-}$ ,  $\text{FeCl}_3^-$ ,  $\text{FeCl}_4^-$  (16). Moreover,  $\text{H}_2\text{SO}_4$  may oxidize  $\text{Fe}^{2+}$  to the potentially immobile species  $\text{Fe}^{3+}$ .

### Selective Separation of $\text{Mg}^{2+}$

#### By Precipitation

The results of selective separation of  $\text{Mg}^{2+}$  from leachate by precipitation are graphically illustrated in Fig. 5. It is noted that  $\text{Mg}^{2+}$  is selectively

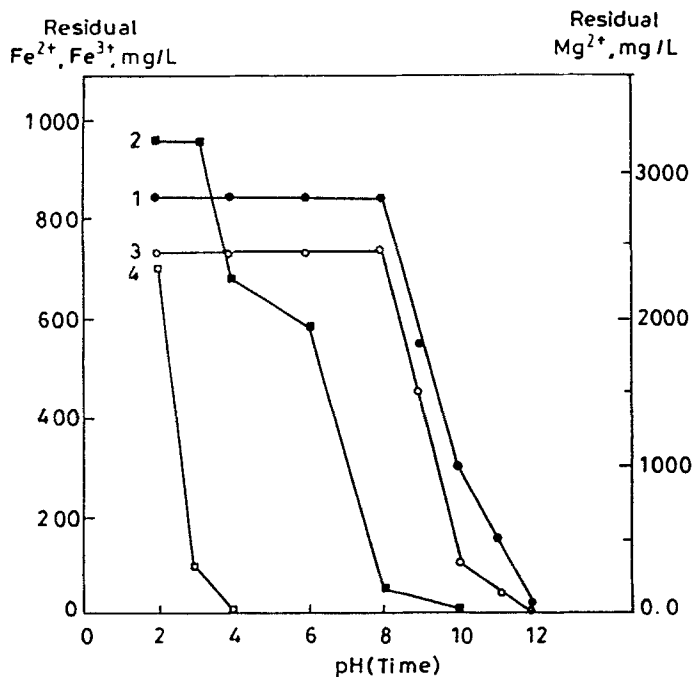


FIG. 5 Selective separation of  $\text{Mg}^{2+}$  by precipitation from leaching solution of olivine (Curves 1 and 2 for  $\text{Mg}^{2+}$  and  $\text{Fe}^{2+}$ , respectively) and serpentinized olivine (Curves 3 and 4 for  $\text{Mg}^{2+}$  and  $\text{Fe}^{3+}$ , respectively) as a function of pH.

and quite effectively separated by precipitation of  $\text{Mg}(\text{OH})_2$  at  $\text{pH} \sim 11.5$ . It is fortuitous that  $\text{Mg}(\text{OH})_2$  precipitates at  $\text{pH} > 10$ , and hence most, if not all, other metal ions are pre-precipitated and separated as hydroxides by filtration. This favors and enhances the precipitation of  $\text{Mg}^{2+}$  as  $\text{Mg}(\text{OH})_2$  without contamination due to coprecipitation.

### By Ionic Flotation

The contaminated metal ions, especially  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$ , are precipitated from the leaching solutions of olivine and serpentinized olivine and filtered. The effect of  $\text{pH}$  on the floatability of  $\text{Mg}^{2+}$  from the filtrate of the leaching solutions was investigated to select the optimum  $\text{pH}$  for maximum flotation efficiency. The results are shown in Fig. 6. As can be seen, the maximum flotation efficiency is obtained in the 7–8.5  $\text{pH}$  range. Separation by flotation is characterized by being less time consuming. Self-flotation of magnesium oleate eliminates the need for air bubbles. It is simple with magnesium oleate to obtain pure magnesium metal.

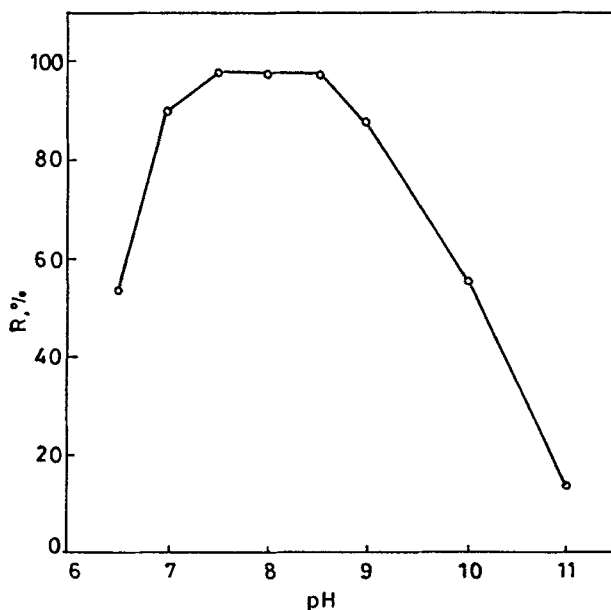


FIG. 6 Ionic flotation of  $\text{Mg}^{2+}$  from the filtrate of the leachate of olivine or serpentinized olivine (containing 2.5 g/L) as a function of  $\text{pH}$ .

## REFERENCES

1. H. Remy, *Treatise on Inorganic Chemistry*, Elsevier, Amsterdam, 1970.
2. H. Vollrath and H. Ingo, *Handbook of Applied Chemistry*, McGraw-Hill, New York, 1983.
3. S. J. Johnstone and M. G. Johnstone, *Minerals for the Chemical and Allied Industries*, 2nd ed., Chapman and Hill, London, 1961.
4. J. A. Dean (Ed.), *Handbook of Chemistry*, 12th ed., McGraw-Hill, New York, 1979.
5. *Mining Chemicals Handbook*, revised ed., American Cyanamid Co., 1986.
6. R. L. Bates and J. A. Jackson (Eds.), *Glossary of Geology*, 2nd ed., American Geological Institute, Falls Church, Virginia, 1980.
7. A. Betekhtin, *A Course of Mineralogy*, Peace Publishers, Moscow, 1969.
8. D. A. Skoog, D. M. West, and F. J. Holler, *Fundamentals of Analytical Chemistry*, 6th ed., Saunders College Publishing, New York, 1992.
9. S. E. Manahan, *Quantitative Chemical Analysis*, Brooks Cole Publishing Co., California, 1986.
10. A. Mizuike, *Enrichment Techniques for Inorganic Trace Analysis*, Springer-Verlag, New York, 1983.
11. S. E. Ghazy, *Sep. Sci. Technol.*, **30**, 933 (1995).
12. A. M. Abeidu and A. B. Moustafa, *J. Mines, Met. Fuels*, **25**, 341 (1977).
13. A. M. Abeidu, *J. Appl. Chem.*, **24**, 245 (1974).
14. F. C. Loughnan, *Chemical Weathering of Silicate Minerals*, American Elsevier, New York, 1969.
15. A. M. Abeidu and A. B. Moustafa, *Indian J. Chem.*, **14A**, 842 (1976).
16. R. B. Fisher and D. G. Petters, *Basic Theory and Quantitative Chemical Analysis*, W. B. Saunders Co., Philadelphia, 1968.

Received by editor July 6, 1995